

## The concept

A chaotic deterministic system  $f_\rho$  has robust extremes under observable  $\phi$  when the associated extreme value statistics depend smoothly on control parameter  $\rho$ .

## Results

Robustness of extremes:

1. depends on system  $f_\rho$  and on observable  $\phi$ ;
2. allows improved estimates by pooling data and
3. improved prediction of (non-stationary) return levels.

## Phenomenology

Robustness of extremes depends on the **system**  $f_\rho$  and on the **observable**  $\phi$ .

Illustration for Lorenz63 model with  $\sigma = 10$ ,  $\beta = \frac{8}{3}$ :

$$\dot{x} = \sigma(y - x), \quad \dot{y} = x(\rho - z) - y, \quad \dot{z} = xy - \beta z. \quad (1)$$

For  $\rho = 28$ : (1) has robust strange attractor [1]. Let

$$\phi_1(x, y, z) = x, \quad \phi_2(x, y, z) = 1 - |x - 5|^{0.25}. \quad (2)$$

Generate time series of length  $10^6$  units (recorded every 0.5) and extract maxima over blocks of 1000 time units. Fit generalised extreme value (GEV) distribution:

$$G(x; \mu, \sigma, \xi) = \exp \left[ - \left( 1 + \xi \frac{x - \mu}{\sigma} \right)_+^{-1/\xi} \right]. \quad (3)$$

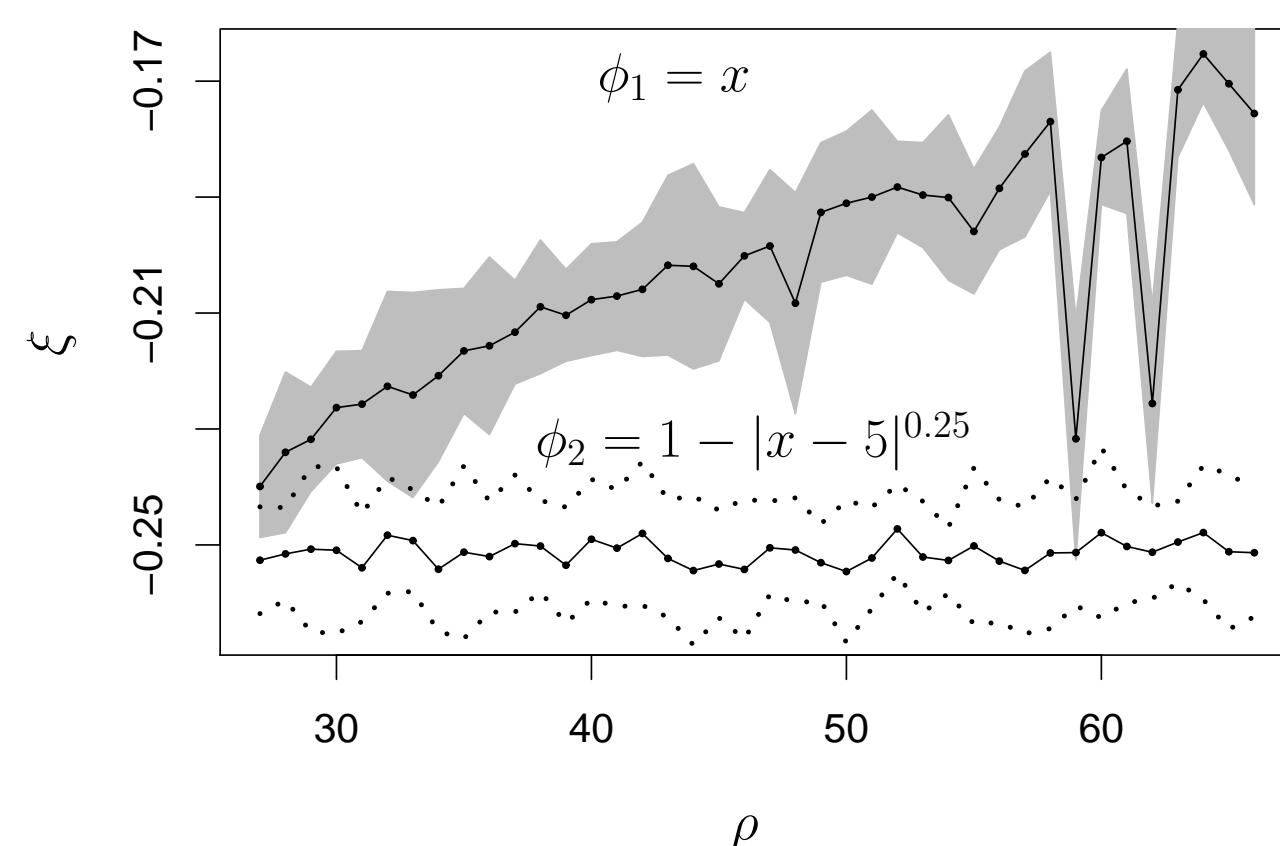


FIG. 1: Maximum likelihood estimates of  $\xi$  for observables  $\phi_1$  and  $\phi_2$  in (2), for  $\rho_j = 27 + j$ ,  $j = 0, 1, \dots$

$\phi_1$ : for small  $\rho$  ( $\approx 28$ ),  $\xi$  varies smoothly with  $\rho$ . Non-linear scaling of attractor  $\rightsquigarrow$  shape of  $\xi(\rho)$ . Discontinuous for large  $\rho$  due to hyperbolicity loss (folds in return map).

$\phi_2$ :  $\xi(\rho) = 0.25$  is constant even under hyperbolicity loss.

**Rigorous proof** available for 1D Lorenz maps.

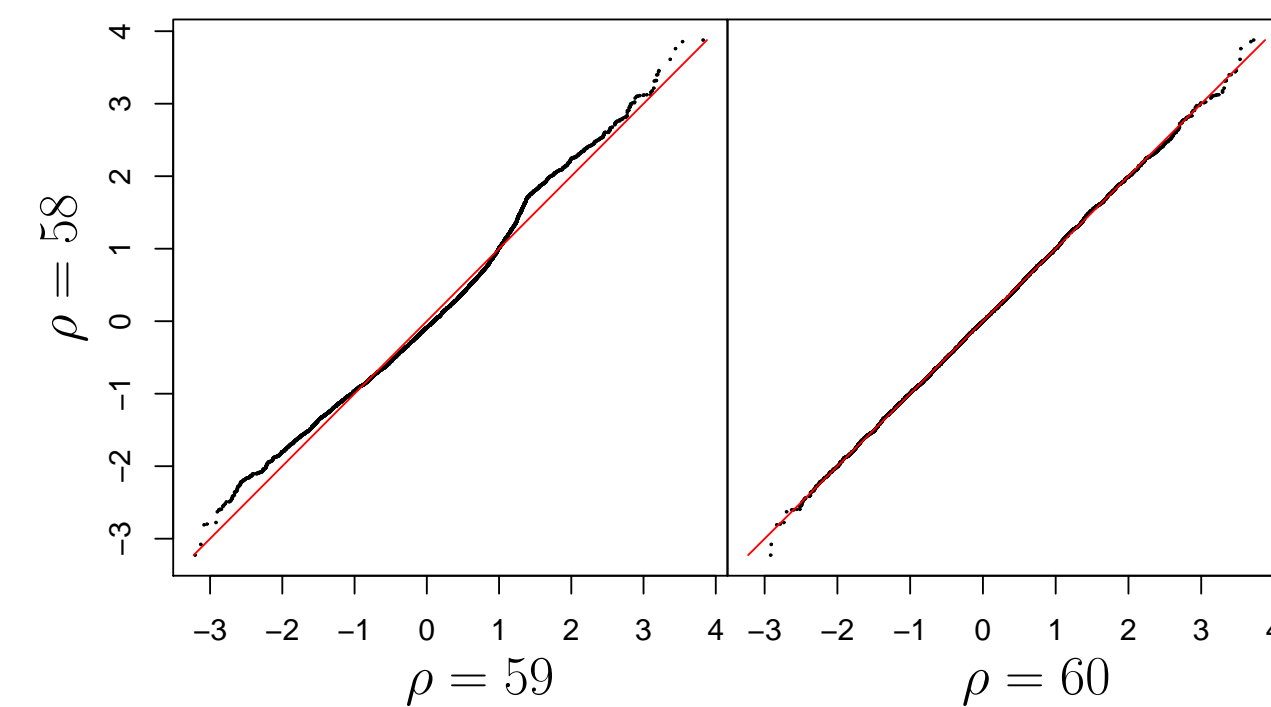


FIG. 2: Discontinuity in upper tail at  $\rho = 59$  for  $\phi_1$  (quantile-quantile plot of  $10^4$  maxima from Fig 1, for each  $\rho = 58, 59, 60$ ).

## Pooling data

Robustness of extremes  $\rightsquigarrow$  enhanced precision of GEV estimators.

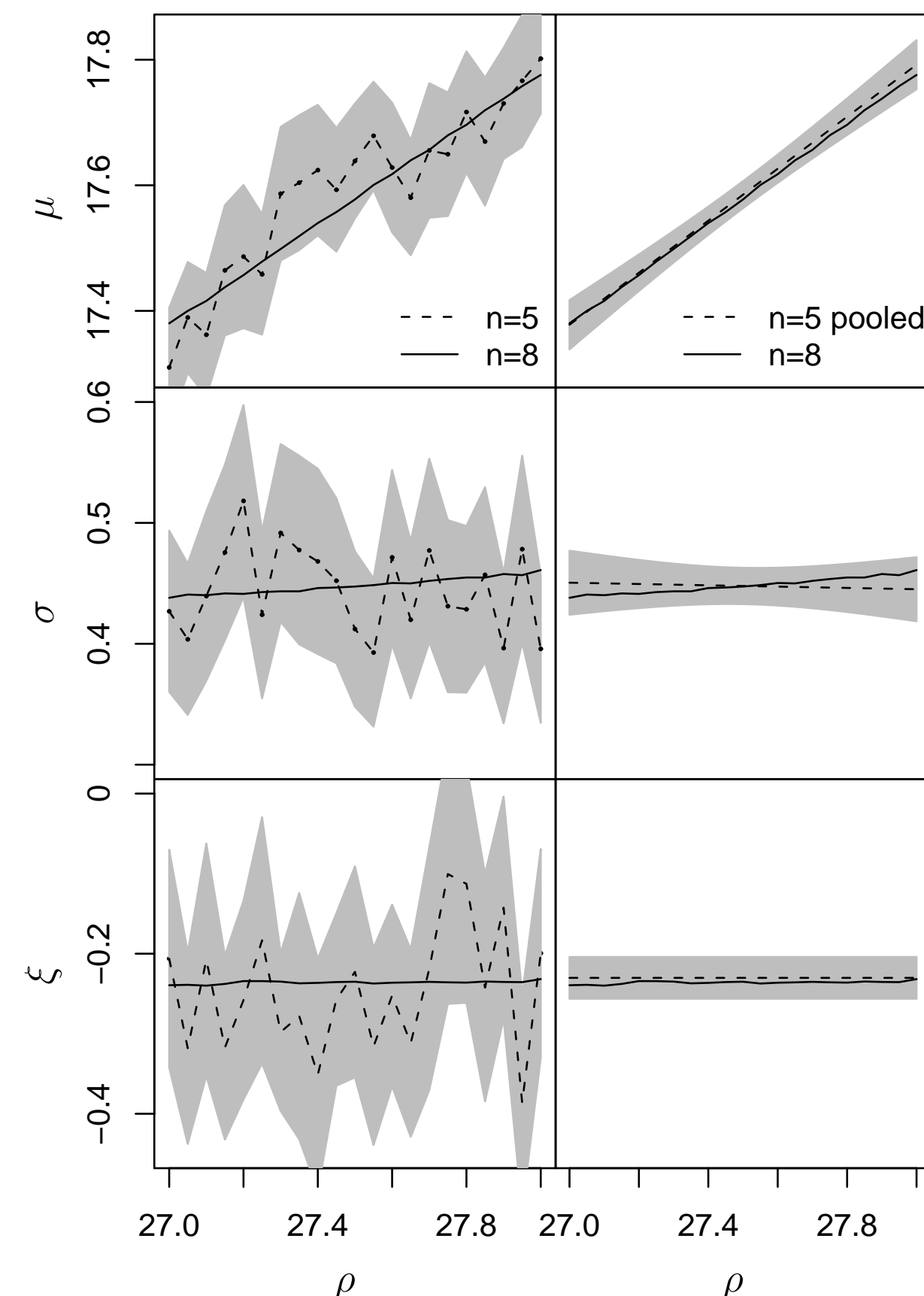


FIG. 3: Left: GEV parameter estimates for (1) with  $n = 5$  (dashed, 95% conf. int. in gray) and  $n = 8$  (solid). Right: “pooled” estimates with  $n = 5$  (dashed, 95% conf. int. in gray).

**Large  $n$ :** GEV parameter estimates  $\rightarrow$  smooth functions of  $\rho$  (solid lines in Fig 3,  $n = 8$ ).

**Small  $n$ :** wild oscillations around “true values” (dashed lines, left column in Fig 3,  $n = 5$ ).

## Enhanced precision: “pooling” short series ( $n = 5$ )

Given robust extremes, information can be pooled from nearby  $\rho$

$\rightsquigarrow$  reduction in uncertainty due to parameter estimation, cfr. grey bands in Fig 3. Assume functional forms

$$\mu(\rho) = \mu_0 + \mu_1 \rho, \quad \sigma(\rho) = \sigma_0 + \sigma_1 \rho, \quad \xi(\rho) = \xi_0.$$

$\xi$  constant in  $\rho$ : approximation, only valid locally. Estimate  $(\mu_0, \mu_1, \sigma_0, \sigma_1, \xi_0)$  by maximum likelihood.

## Prediction & non-stationarity

Robustness of extremes  $\rightsquigarrow$  interpreting and predicting non-stationary extremes.

Robust extreme windspeeds are found in a simple two-layer quasi-geostrophic model [2]:

smooth dependence of windspeed return levels wrt baroclinic forcing parameter  $T_E$  in stationary case.

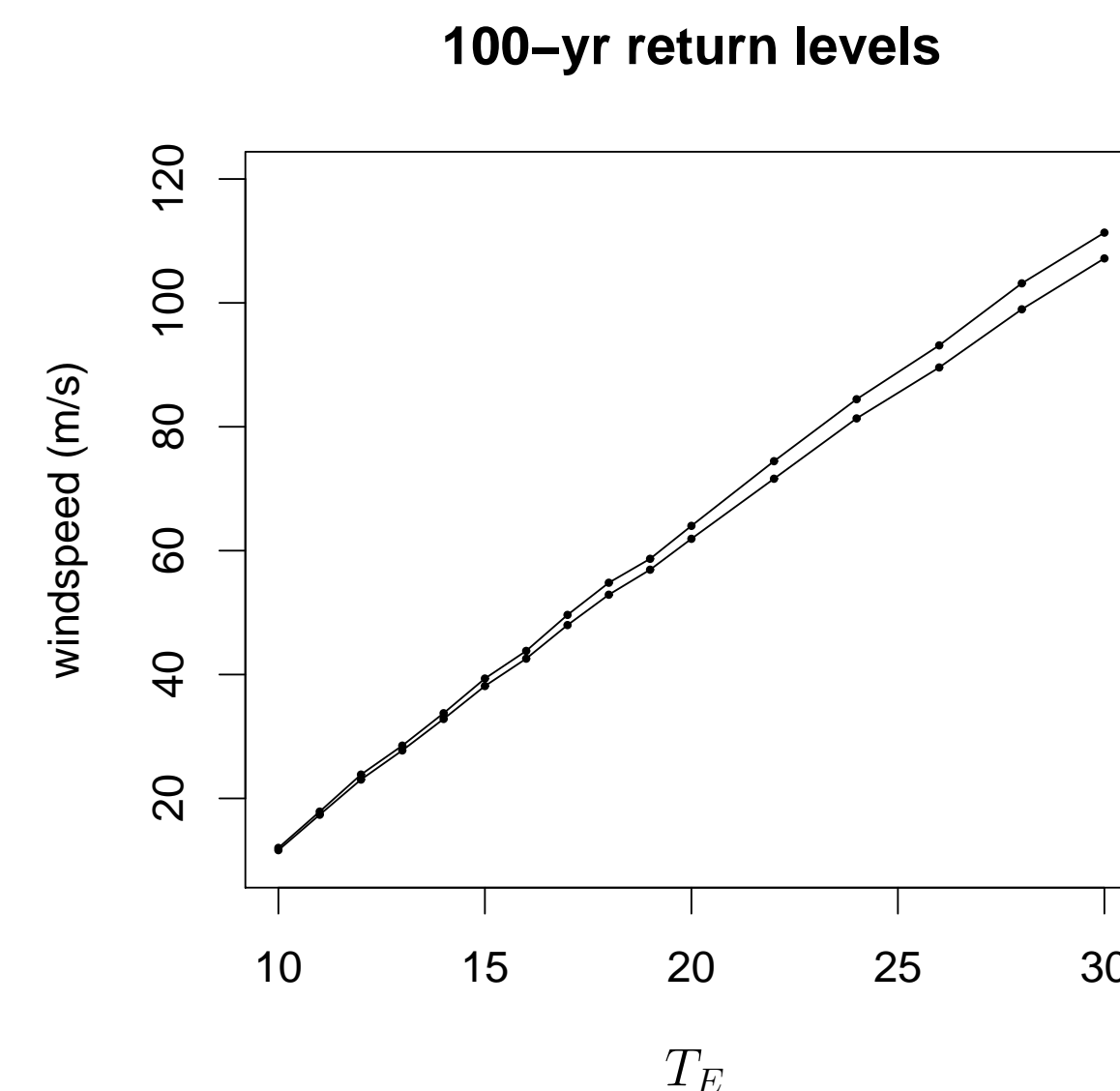


FIG. 4: 100-year windspeed return levels at centre of lower layer, for different values of  $T_E$  (non-pooled GEV fits, stationary case).

Introduce linear time trend in QG model:

$$T_E(t) = (T_E^0 - 1) + t \Delta T_E, \quad t \in [0, t_0], \quad \Delta T_E = 2/300 \text{ yrs.}$$

**Ansatz:** adiabatic + slow trend.

1. trend speed  $\Delta T_E$  is sufficiently small wrt sampling time for upper tail of windspeed distribution;
2. non-stationary extremes remain close (locally in time) to those of stationary system for “frozen”  $T_E(t)$ .

$\Rightarrow$  robustness of extremes wrt **control parameter** translates to smooth change of extremes wrt **time**.

We adopt the Generalized Additive Models for Location, Scale and Shape (GAMLSS) [3]:

1. response distribution is GEV with constant  $\xi$  and cubic smoothing spline for  $(\mu, \sigma)$  with identity link;
2. split sequence of yearly maxima into training and test set (years 1-2250 and 2251-3000);
3. fit non-stationary GEV-GAMLSS to training set;
4. compute time-dependent quantiles and compare to training and test set.

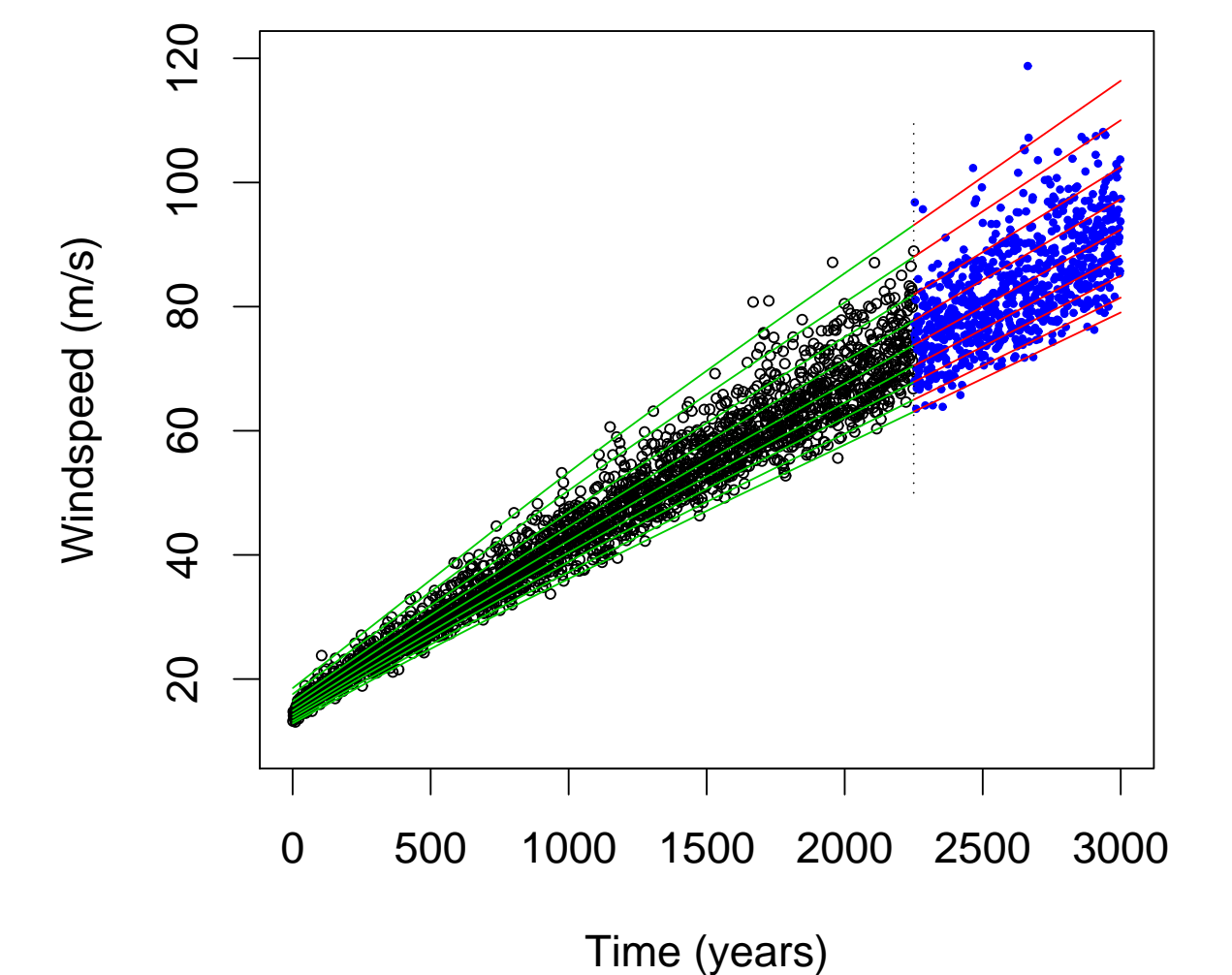


FIG. 5: Points: observed yearly windspeed maxima during training (black) and test (blue) periods. Curves: time-dependent estimated quantiles from GEV-GAMLSS.

quantiles	0.4	2	10	25	50	75	90	98	99.6
training	0.5	2.3	9.5	25.2	50.3	76.1	90.0	98.0	99.6
test	0.7	3.6	14.3	30.5	51.7	74.7	91.3	97.7	99.3

Fraction of points below the estimated quantile curves in Fig. 5 (corresp. to top row) during training (green, centre row) and test (red, bottom row) periods.

Illustrates **potential for predicting return levels in a non-stationary system exhibiting robust extremes**.

## References

- [1] Morales, C. A., Pacifico, M. J., and Pujals, E. R. *Proc. Amer. Math. Soc.* **127**(11), 3393–3401 (1999).
- [2] Felici, M., Lucarini, V., Speranza, A., and Vitolo, R. *J. Atmos. Sci.* **64**(7), 2137–2158 Jul (2007).
- [3] Rigby, R. A. and Stasinopoulos, D. M. *J. Roy. Statist. Soc. Ser. C* **54**(3), 507–554 (2005).